

# Integrating the Analytical Hierarchy Process (AHP) and the Frequency Ratio (FR) Model in Landslide Susceptibility Mapping of Shiv-khola Watershed, Darjeeling Himalaya

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**Abstract** To prepare a landslide susceptibility map of Shiv-khola watershed, one of the landslide prone parts of Darjeeling Himalaya, remote sensing and GIS tools were used to integrate 10 landslide triggering parameters: lithology, slope angle, slope aspect, slope curvature, drainage density, upslope contributing area (UCA), lineament, settlement density, road contributing area (RCA), and land use and land cover (LULC). The Analytical Hierarchy Process (AHP) was applied to derive factor weights using MATLAB with reasonable consistency ratio (CR). The frequency ratio (FR) model was used to derive class frequency ratio or class weights that indicate the relative importance of individual classes for each factor. The weighted linear combination (WLC) method was used to determine the landslide susceptibility index value (LSIV) on a GIS platform, by incorporating both factor weights and class weights. The Shiv-khola watershed is classified into five landslide susceptibility zones. The overall classification accuracy is 99.22 and Kappa Statistics is 0.894.

**Keywords** Analytical Hierarchy Process (AHP), frequency ratio (FR) model, India, landslide susceptibility, West Bengal

## 1 Introduction

The identification of the causative factors is the basis of many methods of landslide susceptibility assessment. The spatial distribution of slope instability is essential for land use planning. Landslides are the result of two interacting sets of forces: (1) natural precondition factors that govern the stability conditions of slopes; and (2) preparatory and triggering factors caused by either natural factors or human intervention. Landslide analysis is mainly done by assessing susceptibility, hazard, and risk. The remote sensing and GIS based landslide hazard zonation approach was studied by Nautiyyal (1966), Muthu and Petrou (2007), Wu and Qiao (2009), and other researchers. Rowbotham and Dudycha (1998), Donati and Turrini (2002), Lee and Choi (2003), Lee et al. (2004), Lee, Choi, and Min (2004), Lee and Pradhan (2007), Pradhan

and Lee (2010a, 2010b), Sarkar and Kanungo (2004), and Pandey et al. (2008) studied and applied the probabilistic model for landslide susceptibility and risk evaluation. Guzzetti et al. (1999), Pistocchi, Luzi, and Napolitano (2002) and Dai and Lee (2002) summarized many landslide hazard evaluation studies. Jibson, Edwin, and John (2000) and Zhou et al. (2002) applied the probabilistic models for landslide risk and hazard analysis. Vijith and Madhu (2008) introduced the logistic regression model for landslide hazard mapping.

The models in connection to slope stability and shallow and deep seated landslides were introduced and verified by Bhattarai and Aoyama (2001) and Bradinoni and Church (2004). The geotectonic factors of slope instability were studied in detail by Carson (1975), Windisch (1991), and Borga et al. (1998). Comprehensive lists of stability factors commonly employed in the factors mapping approach were given by Guzzetti et al. (1999) and Tiwari and Marui (2003, 2004).

The Analytical Hierarchy Process (AHP), a semi-quantitative method based on decomposition, comparative judgment, and synthesis of priorities is often useful for regional susceptibility studies as suggested by Saaty (1980), Yalcin and Bulut (2007), and Yalcin (2008). The frequency ratio (FR) model has also become a popular and realistic quantitative approach in landslide susceptibility mapping. This approach is related to historical landslide events and their areal coverage. Lee and Pradhan (2007) argued that the frequency ratio model provides a correlation between historical slide locations and various influencing factors under consideration. Intarawichian and Dasananda (2011) applied the frequency ratio model to analyze slope instability and treated the model as a popular quantitative method.

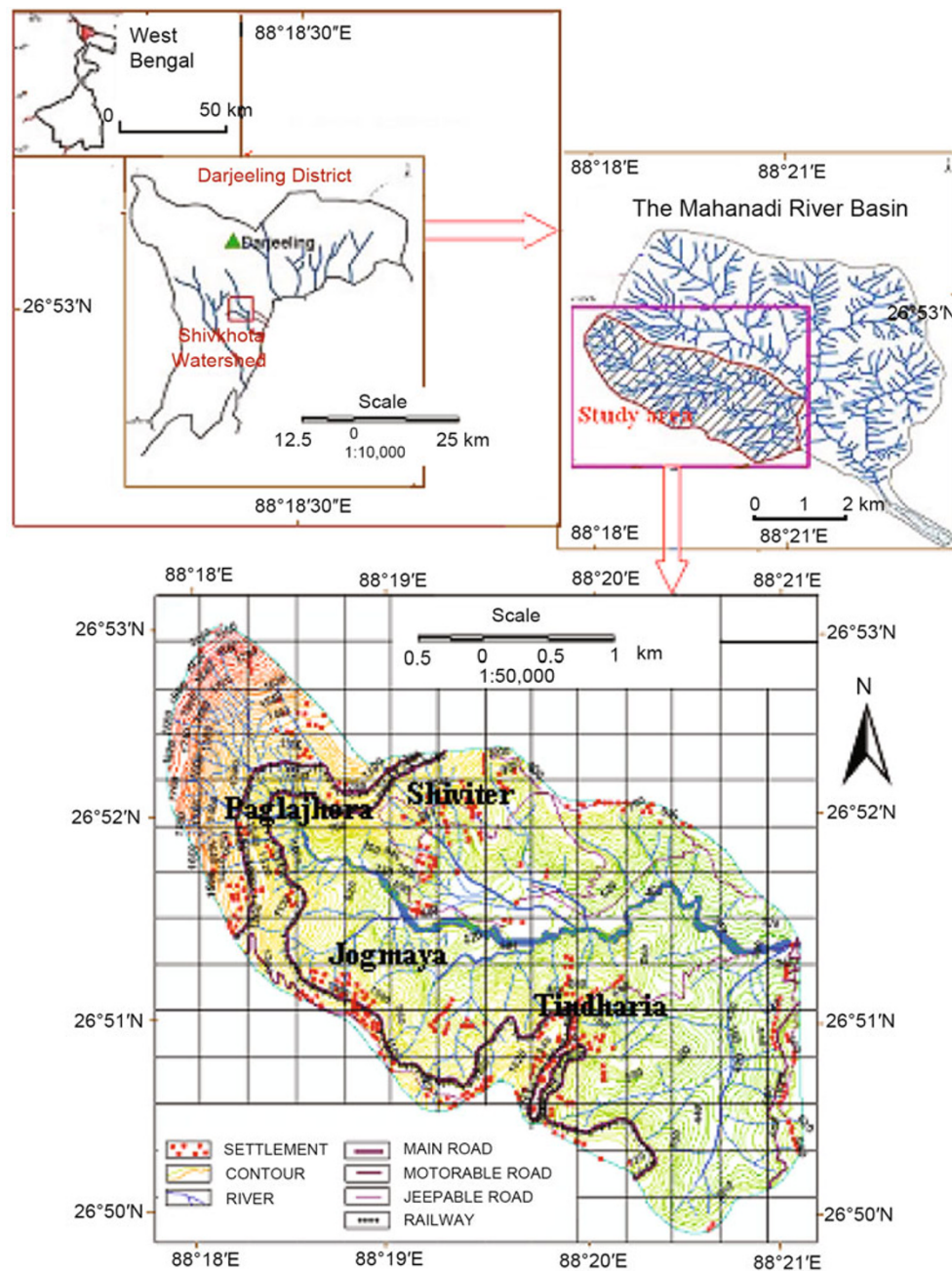
This study deals with the estimation of factor weights and class frequency ratios using the AHP and FR model respectively. Integration between factor weight (FW) and class frequency ratio (FR) was performed with the help of a linear combination model. This is done to derive pixelwise landslide susceptibility index values (LSIV) and prepare a landslide susceptibility map.

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Tectono-statigraphically, the study area, Shiv-khola watershed (Figure 1) is located in the southern escarpment slope of Darjeeling Himalaya, where high-grade metamorphic rocks of the Darjeeling and Chungthang groups are thrust over low-grade metamorphic rocks of the Daling Group along the Main Central Thrust (MCT) (Mallet 1875; Sinha-Roy 1982). The MCT and the Main Boundary Thrust (MBT) pass through the study area (Figure 2). The MCT (a major ductile shear zone) divides two major litho-tectonic units, the Higher Himalayan Crystalline Sequence (HHCS) and the Lesser Himalayan Sequence (LHS) in Darjeeling Himalaya. The HHCS

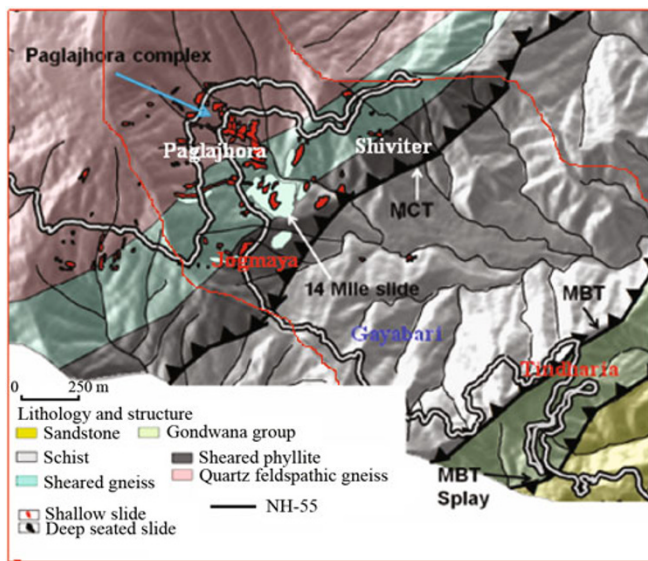
comprises quartzo-feldspathic gneisses of both igneous and sedimentary origin that have been subjected to a high grade of metamorphism (Catlos et al. 2001). The LHS is dominated by garnet-biotite-mica schist and chlorite schists in the upper part and slates and phyllites in the lower part. The landslide affected areas are Paglajhora, Tindharia, Gayabari, Mahanadi, Jogmaya and Shiviter. During the rainy season water percolates through the exposed rock joints and entrains the finer particles and reduces the cohesive strength of the soil.

Rapid urbanization and expansion of tourism in Darjeeling Himalaya are putting unprecedented pressure on land and



**Figure 1.** Location map of the Shiv-khola watershed in West Bengal, India

Source: Mondal and Maiti (2011).



**Figure 2. Tectono-stratigraphy and past landslides in the study area**

Source: Geological Survey of India (East Kolkata), 2009.

soil with the gradual elimination of virgin forest land after independence. Lack of land use planning coupled with vulnerable geological structures and frequent heavy rainfall have led to the formation of a vicious cycle of soil erosion and landslides during and after the monsoon seasons, causing devastating damage to human lives and properties. Significant studies in the Darjeeling Himalaya identified the causes and consequences of major landslide occurrences (Dutta 1966). Since 1968 the Shiv-khola watershed has experienced 128 reachable landslide events, of which 76 were considered as reactivated (less than 70 m from old landslides) and 52 as fresh events (70 m or more from old landslides). These landslide events took place in 16 years, out of which 12 were considered major landslide years. All the landslide events occurred during the monsoon period with continuous and heavy showers. Rainfall on all the major landslide event dates was more than the critical rainfall calculated after Borga et al. (1998). Most of the landslide events occurred in the lithological unit of Darjeeling Gneiss, Daling, Damuda, and Siwalik.

In the Shiv-khola watershed, Lower Paglajhora, Tindharia, Shiviter, Gayabari, and Mahanadi are the major and prominent landslide locations where settlements, communication lines, and tea garden areas are being affected severely by the frequent occurrence of landslides. Since 1968, Paglajhora alone has had 10 landslide events, all in the above-mentioned landslide event years. The majority of these landslides was dangerous as in most of the events Hill Cart Road (NH-55) was affected and the communication line between Siliguri and Darjeeling was completely interrupted, from days to months. Paglajhora sinking zone faced massive slope failures in 1998, 2002, 2005, and 2011, which indicates that the occurrence of landslides in the region is ongoing. This poses a

tremendous threat to upslope settlements and the Hill Cart Road (life line between Siliguri and Darjeeling Town). The landslide events at Tindharia also frequently cut off the Hill Cart Road and threaten the safety of tourists, upslope settlements, and tea gardens. In Shiviter, around eight acres of land were destroyed by destructive slope failure in the past 10 years. The physiographic configuration (arcuate) that provides a favorable condition for producing hydrostatic pressure, proximity to the Main Central Thrust (MCT) and the Main Boundary thrust (MBT), intensely fractured and sheared bedrock, toe cutting and headward erosion of debris covered slope by first-order tributaries, immense pressure over the fragile slope materials from manmade concrete structures, moderate to steep slope gradient, improper drainage network orientation, and accumulation of highly anisotropic materials with a great thickness and low shearing resistance have made these landslide locations in the Shiv-khola watershed most unstable in character.

## 2 Data and Methods

In this study, thematic data layers of all the landslide inducing factors were integrated to prepare a landslide susceptibility map using a linear combination model in GIS. The Analytical Hierarchy Process (AHP) was used to derive the prioritized factor rating value (PFRV) and a Frequency Ratio (FR) model was applied to obtain the prioritized class rating value (PCRv) for all the landslide triggering factors considered in the study. The integration between PFRV and PCRv was made in a linear combination model on a GIS platform to estimate the landslide susceptibility index value (LSIV) for each pixel and a suitable classification technique was incorporated to prepare the landslide susceptibility map of the Shiv-khola watershed. The data used in the study are: satellite image (IIRS P6/Sensor-LISS- III, Path-107, Row-052, date 18 March 2010); modified shuttle radar topography mission (SRTM) data with scene size 1 degree latitude and 1 degree longitude (date 5 April 2008); Google Earth image (1 September 2010); geological map (Geological Survey of India, East Kolkata); and topographic map (78B/5, Survey of India). Data layers for landslide inducing factors were generated using ERDAS Imagine 8.5, ArcView, and ArcGIS Software.

## 3 Analyses

The following section presents the methods and results of the landslide analyses in this study.

### 3.1 Determination of Landslide Triggering Factors

The landslide triggering factors were identified by interviewing the local people and an investigation of the landslide sites in the watershed through intensive fieldwork. During the 10 days fieldwork in July 2011, landslide locations were



identified with GPS, and lithological structure, land use and land cover type around the landslide scars, slope angles, construction of manmade structures and their role in promoting landslides, drainage networks, altitude, and slope aspects were investigated to determine important landslide triggering factors. Ten landslide triggering factors including lithology, slope angle, drainage, slope aspect, slope curvature, lineament, upslope contributing area (UCA), land use and land cover (LULC), road contributing area (RCA), and settlement density were taken into account to prepare the landslide susceptibility map of the Shiv-khola watershed and their hierarchical arrangement was made on priority basis. Shiv-khola watershed is a small mountain basin where rainfall is uniformly distributed over the entire area, so rainfall intensity was not considered in the landslide susceptibility calculation (Mondal and Maiti 2011).

### 3.2 Generation of Landslide Inducing Factor Maps

First, the contour map at 20 m intervals was digitized from the Survey of India (SOI) topographic map (1987, 78B/5) at the scale of 1:50,000 and subsequently employed for generating the digital elevation model (DEM) using the ArcGIS Software. Then slope gradient, slope curvature, and slope aspect maps were derived from DEM with 25 m × 25 m grid cell size and a supervised classification was made to derive all these parameters in raster value domain following the earlier works of Dhakal, Amada, and Aniya (2000). Surface curvature is a topographic attribute that describes the convexity/concavity of a terrain surface. Curvature depicts the slope gradient or slope direction (aspect), usually in a particular direction (Gallant and Wilson 2000). A positive curvature indicates the surface is upwardly convex at a grid cell and a negative curvature indicates the surface is upwardly concave at that grid cell. A value of zero indicates the surface is flat. The expected values of all three output raster images for a hilly area can vary from −0.5 to 0.5; for steep, rugged mountains the value can vary between −4 and 4.

The lithological map of the study area was collected from the Geological Survey of India (GSI), Kolkata (Eastern Region) and necessary modifications were incorporated after intensive field investigation. The final lithological map was made with seven rock types and transformed into raster value domain in ArcGIS. Class weight value for each lithological class was assigned according to rock mass strength, described by GSI. A drainage density map (length of drainage/m<sup>2</sup>) was made at the grid resolution of 23.5 m × 23.5 m from the topographic map (78B/5) and classified into 10 equal intervals.

Lineament indicates the zone of weakness, representing some linear to curvilinear features such as fracture, joint, and fault in the geological structure. There is no basic difference between these three features. All these linear to curvilinear features were identified as the same deformed surface where the propensity of slope instability is very high. To generate a lineament map (distance from lineament in meters) of the

Shiv-khola watershed, PCI-GEOMATICA was used and in the extraction process three SRTM bands were taken into account: Near Infrared (Band-I, 0.7–1.3 μm), Red (Band-II, 0.6–0.7 μm), and Green (Band-III, 0.5–0.6 μm). The algorithm used to prepare the lineament map is Lineament Extraction. The study area was classified into 10 classes on the basis of distance (m) from lineaments.

Upslope contributing area (UCA) is an effective indicator of drainage concentration over space. The place with more contributing area encompasses more soil saturation and reduces soil cohesion. Specific contributing area (total contributing area divided by the contour length) is computed by distributing flow from a pixel among its entire lower elevation neighbor pixel (Borga et al. 1998) (Eq. 1). An upslope contributing area map was prepared based on the calculated contributing area value for each 0.25 km<sup>2</sup> grid and the map was divided into 6 equal classes.

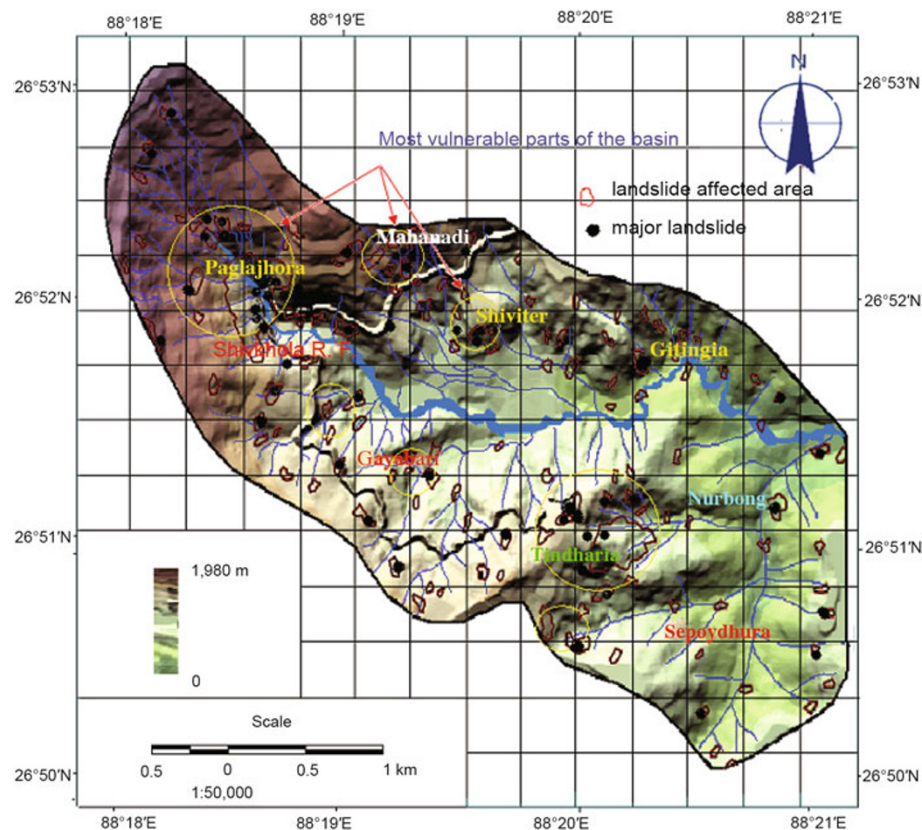
$$F_i = \frac{S_i L_i}{\sum_i S_i L_i} \quad \text{Eq. 1}$$

where, the summation ( $\sum_i$ ) is for the entire lower neighbors,  $S$  is the directional slope, and  $L$  is the effective contour length that acts as the weighting factor. The value of  $L$  used here is 10 m (the pixel size) for the cardinal neighbors and 14.14 m (the pixel diagonal) for diagonal neighbors.

The road contributing area (RCA) map was made by multiplying road contributing length (RCL) with road contributing width (RCW) and was classified into eight equal classes from the concerned topographic map and converted into raster value domain in ArcGIS. The settlement density map was prepared by applying a 3×3 kernel in ArcGIS and the basin was classified into seven equal density classes. The land use and land cover (LULC) map of the watershed was prepared with the help of the LISS-III satellite image (2010) and the Google Earth image in consultation with the SOI topographic map (78B/5). After verifying the ground truth with GPS a land use and land cover map was developed in GIS. The Shiv-khola watershed was classified into 10 individual land use types (bare surface, agricultural land, jungle, roads, settlement, tea garden, open forest, degraded forest, mixed forest, and dense forest).

### 3.3 Landslide Inventory Map

A landslide distribution or inventory map (Figure 3) was created to determine landslide affected areas (%) and frequency of landslides for each class of the landslide inducing factors. Mondal and Maiti (2011) identified major and minor landslide locations during field investigation and mapped them by evaluating the SOI topographic map (78B/5), satellite image (IRS LISS- III, 2000), SRTM data (2008), and Google Earth image (2000). Several field investigations were conducted to identify the landslide locations as well as to cross-check the prepared landslide map. Then, the map was digitized and converted into raster value domain in ArcGIS. All the landslide



**Figure 3. Landslide inventory map of the Shiv-khola watershed in West Bengal, India**

Source: Mondal and Maiti (2011).

triggering factor maps were incorporated with this landslide inventory map to understand the degree of importance of each factor in landsliding.

### 3.4 The Analytical Hierarchy Process (AHP) and the Prioritized Factor Rating Value (PFRV)

AHP is a decision-making and semi-quantitative value judgment approach that serves the objectives of the decision makers. This process is employed in this study to support the decision on the instability rank of the factors by estimating the prioritized factor rating value (PFRV). In the AHP, different factor preferences and their conversion into numerical values were accomplished with the help of comparative oral judgment based on interviewing the local people inhabiting in the landslide prone area and synthesis of priorities (Table 1). A pairwise comparison matrix for the study area was constructed on the basis of the preference of a factor as compared with the other factor and arithmetic mean method was applied to arrange landslide triggering factors hierarchically and to determine the prioritized factor rating value/eigenvector (PFRV) with reasonable consistency ratio (CR), based on Saaty (1977, 1980) and Saaty and Vargas (2000), using MATLAB (Table 2). To develop the pairwise comparison matrix, each factor was rated against every other factor by assigning a relative dominant value ranging between 1 and 9 on the basis of the relative importance of the factors in

relation to landslide frequency. The value also varies between the reciprocals 1/2 and 1/9 for inverse comparison (Table 1).

Another appealing feature of the AHP is the ability to evaluate pairwise rating inconsistency. The eigenvalues enable the quantification of a consistency measure that is an indicator of the inconsistencies or intransitivities in a set of pairwise ratings. Saaty and Vargas (2000) stated that for a consistent reciprocal matrix, the largest eigenvalue  $\lambda_{\max}$  is equal to the number of comparisons  $n$ . An index of consistency, known as the CR (Consistency Ratio), is used to indicate the probability that the matrix judgments were randomly generated (Saaty 1977).

$$CR = CI / RI \quad \text{Eq. 2}$$

where RI is the average of the resulting consistency index depending on the order of the matrix given by Saaty and CI is the consistency index that is expressed in Eq. 3. If the value of CR is smaller or equal to 10 percent, the inconsistency is acceptable, but if the CR is greater than 10 percent, the subjective judgment needs to be revised (Saaty 1977).

$$CI = \lambda_{\max} - n / n - 1 \quad \text{Eq. 3}$$

Saaty and Vargas (2000) randomly produced reciprocal matrices using scales 1/9, 1/8, 1/7...1...8, 9 to evaluate a so-called random consistency index (RI). The average RI of 500 matrices is given in Table 3.

**Table 1. Scale of preference between two parameters**

Scale	Degree of Preference	Explanation
1	Equally	Two activities contribute equally to the objective
3	Moderately	Experience and judgment slightly to moderately favor one activity over another
5	Strongly	Experience and judgment strongly or essentially favor one activity over another
7	Very Strongly	An activity is strongly favored over another and its dominance is showed in practice
9	Extremely	The evidence of favoring one activity over another is of the highest degree possible of an affirmation
2, 4, 6, and 8	Intermediate values	Used to represent compromises between the references in weight 1, 3, 5, 7, and 9
Reciprocals	Opposites	Used for inverse comparison

Source: Saaty and Vargas (2000).

**Table 2. Landslide triggering factors and prioritized factor rating values (weights) in the Shiv-khola watershed, West Bengal, India**

Factors	1	2	3	4	5	6	7	8	9	10	Prioritized Rating (PFRV)
(1) Slope	1	2	3	4	5	6	7	8	9	9	0.2944
(2) Lithology	1/2	1	2	3	4	5	6	7	8	9	0.2150
(3) Drainage	1/3	1/2	1	2	3	4	5	6	7	8	0.1537
(4) Lineament	1/4	1/3	1/2	1	2	3	4	5	6	7	0.1087
(5) Curvature	1/5	1/4	1/3	1/2	1	2	3	4	5	6	0.0764
(6) UCA	1/6	1/5	1/4	1/3	1/2	1	2	3	4	5	0.0535
(7) RCA	1/7	1/6	1/5	1/4	1/3	1/2	1	2	3	4	0.0375
(8) LULC	1/8	1/7	1/6	1/5	1/4	1/3	1/2	1	2	3	0.0266
(9) Settlement Density	1/9	1/8	1/7	1/6	1/5	1/4	1/3	1/2	1	2	0.0193
(10) Slope Aspect	1/9	1/9	1/8	1/7	1/6	1/5	1/4	1/3	1/2	1	0.0149

CI (consistency index) = 0.0615; RI (random consistency index) = 1.49; and CR = 0.0413 (consistent)

RCA = Road contributing area; UCA = Upslope contributing area; LULC = land use and land cover

**Table 3. Random consistency index (RI)**

N	1	2	3	4	5	6	7	8	9	10	11	12	13	14	15
RI	0	0	0.58	0.90	1.12	1.24	1.32	1.41	1.45	1.49	1.51	1.53	1.56	1.57	1.59

Source: Saaty (1977).

### 3.5 Frequency Ratio (FR) Model and Prioritized Class Rating Value (PCRVR)

The frequency ratio (FR) model is a well accepted and popular quantitative approach for the preparation of landslide susceptibility maps. Lee and Talib (2005), Lee and Pradhan (2007), Jadda (2009), Avinash and Ashamanjari (2010), and Intarawichian and Dasananda (2011) successfully applied the FR model to generate landslide susceptibility zoning maps. To obtain the frequency ratio (FR) for each class of all the data layers a combination has been established between the landslide inventory map and factor maps using the following equation.

$$FR = \frac{N_{\text{pix}(S_i)} / N_{\text{pix}(N_i)}}{\sum_i N_{\text{pix}(S_i)} / \sum_i N_{\text{pix}(N_i)}} \quad \text{Eq. 4}$$

where  $N_{\text{pix}(S_i)}$  is the number of pixels containing slide in class  $i$ ,  $N_{\text{pix}(N_i)}$  is the total number of pixels having class  $i$  in

the watershed,  $\sum_i N_{\text{pix}(S_i)}$  is the total number of pixels containing landslide,  $\sum_i N_{\text{pix}(N_i)}$  is the total number of pixels in the watershed.

The derived frequency ratio (FR) value of more than 1 indicates strong and positive relationship between landslide occurrences and the concerned class of the data layer and high landslide susceptibility, whereas a FR value of less than 1 depicts negative relationship and low landslide susceptibility. In this study, the FR value for each class is accepted as prioritized class rating value or prioritized class weight.

### 3.6 Linear Combination Model and Landslide Susceptibility Classification

Avinash and Ashamanjari (2010) and Intarawichian and Dasananda (2011) used a landslide susceptibility index value (LSIV), which is the summation of class- and factor-weighted values. FR values for each class (PCRVR or prioritized class

weight) as well as prioritized factor rating values (PFRV) for each factor map were taken into account in calculating the landslide susceptibility index value (LSIV) with the following linear combination model:

$$\text{LSIV} = \sum_{i=1}^n (W_i \times \text{FR}_i) \times \text{FV} \quad \text{Eq. 5}$$

where,  $n$  is the total number of factors included in the study ( $n = 10$ ),  $W_i$  is factor weight (PFRV), FV is factor value, and  $\text{FR}_i$  is class frequency ratio or prioritized class weight.

In this study, the LSIV varied from 4.81 to 16.00. The higher the value of LSIV, the greater was the propensity of landslide occurrence and vice versa. The frequency distribution of landslide susceptibility index values shows that the LSIV based frequency curve has many oscillations. To produce a better classification of the watershed into landslide susceptibility zones, moving average with averaging window lengths of 3, 5, 7, and 9 was considered for smoothing the frequency distribution curve (Figure 4). After analyzing the four new curves, the watershed was classified into five landslide susceptibility zones: Very Low, Low, Moderate, High, and Very High, with class boundaries at the significant changes of gradient of these curves. The abrupt change points on frequency curves (landslide threshold boundaries) were 7.05, 9.29, 11.5, and 13.8. A  $3 \times 3$  "majority filter" was applied to the map as a post-classification filter to reduce the high frequency variation.

To verify the landslide susceptibility map, landslide density under each susceptibility class was computed. The landslide inventory map was crossed with the calculated landslide susceptibility map to derive landslide affected pixels for each susceptibility class (zone). Research by Sarkar and Kanungo (2004) indicates that the higher the landslide density, the greater is the probability of and the area affected by landslides in a landslide susceptibility class.

### 3.7 Accuracy Assessment of the Landslide Susceptibility Map with Field Data

Accuracy was assessed by comparing the classification with geographical data that are assumed to be true using Erdas Imagine (8.5). Ground truth verification data were obtained with the help of GPS from 50 existing landslide locations. Simultaneously, 50 randomly selected reference pixels from the classified image corresponding to the 50 landslide locations (GPS record) were used for evaluating the validity of the landslide susceptibility map (Congalton 1991).

## 4 Landslide Susceptibility of the Shiv-khola Watershed

An effective management to prevent slope failure deals with the triggering factors and their roles in landsliding. The following section presents the relationship between various landslide inducing factors and landslide susceptibility, as well as landslide susceptibility characterization in the study area.

### 4.1 The Relationship between Landslide Susceptibility and Triggering Factors

Landslide susceptibility of the Shiv-khola watershed was affected by the interaction between landslide triggering factors and existing landslides. Class frequency ratio indicates the relative importance of individual classes for each factor and provides important information for analyzing the role of these factors in inducing landslides. The class frequency ratios (prioritized class weights) of the 10 landslide triggering factors are presented in Table 4. (1) Slope gradient of the watershed varies from very gentle (around  $10^\circ$ ) in the

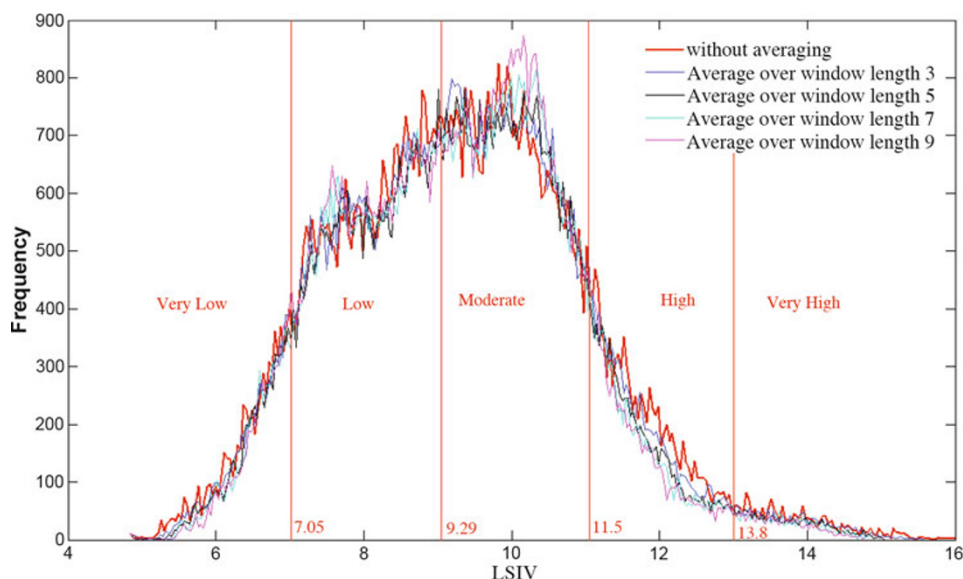


Figure 4. Frequency distribution of landslide susceptibility index value of the Shiv-khola watershed in West Bengal, India



**Table 4. Class frequency ratio (prioritized class rating value)**

Slope Gradient (degree)-1					
Classes	Number of Pixels [ $N_{pix(N_i)}$ ]	% of $N_{pix(N_i)}$	Landslide Pixels [ $N_{pix(S_i)}$ ]	% of $N_{pix(S_i)}$	FR/PCRV
0–7.17	3,353	10.12	190	5.63	0.53
7.17–14.34	3,238	9.77	202	5.99	0.61
14.34–19.92	3,587	10.83	211	6.26	0.58
19.92–24.97	2,445	7.38	201	5.96	0.81
24.97–29.75	3,555	10.73	311	9.22	0.86
29.75–34.53	2,776	8.38	329	9.75	1.16
34.53–39.57	3,854	11.63	413	12.24	1.05
39.57–45.95	3,276	9.89	417	12.36	1.25
45.95–54.71	3,557	10.74	523	15.51	1.44
54.71–67.73	3,490	10.53	626	18.56	1.76
Slope Aspect (direction of slope)-2					
Flat	784	2.37	24	0.711	0.30
North	3,879	11.71	665	19.72	1.68
Northeast	3,797	11.46	443	13.13	1.15
East	4,346	13.12	675	20.01	1.53
Southeast	6,290	18.99	789	23.39	1.23
South	4,556	13.75	597	17.70	1.29
Southwest	3,332	10.06	35	0.74	0.07
West	2,870	8.66	69	2.05	0.24
Northwest	3,277	9.89	76	2.25	0.23
Slope Curvature (positive, negative, and zero)-3					
–25.87– –11.41	995	3.00	221	6.55	2.18
–11.41– –5.73	785	2.37	210	6.23	2.63
–5.73– –2.33	2,111	6.37	486	14.40	2.26
–2.33– –0.63	2,431	7.34	374	11.09	1.51
–0.63–0.50	10,045	30.32	388	11.50	0.38
0.50–2.49	6,302	19.02	268	7.95	0.42
2.49–7.31	5,438	16.41	464	13.76	0.84
7.31–14.69	3,343	10.09	475	14.08	1.40
14.69–24.33	895	2.70	222	6.58	2.44
24.33–46.45	786	2.37	265	7.86	3.32
Lineaments (distance from lineament, m)-4					
0–57.42	3,381	10.20	624	18.50	1.81
57.42–126.32	3,786	11.43	668	19.80	1.73
126.32–229.68	3,695	11.15	451	13.37	1.20
229.68–356.00	3,252	9.82	522	15.48	1.58
356.00–528.26	4,799	14.48	444	13.16	0.91
528.26–723.45	4,141	12.50	286	8.48	0.68
723.45–964.65	3,887	11.73	221	6.55	0.56
964.65–1251.75	3,921	11.83	120	3.56	0.30
1251.75–1642.20	1,419	4.29	37	1.10	0.26
1642.20–2925.40	850	2.57	0	0	0
Drainage Density (length of drainage, km/km <sup>2</sup> )-5					
0–1.90	5,560	16.78	90	2.67	0.16
1.90–3.80	5,453	16.46	109	3.23	0.20
3.80–5.71	3,289	9.93	158	4.68	0.47
5.71–7.61	5,049	15.24	137	4.06	0.27
7.61–9.51	3,477	10.49	159	4.71	0.45
9.51–11.41	2,728	8.23	455	13.49	1.64
11.41–13.31	1,875	5.66	534	15.83	2.80
13.31–15.21	2,191	6.61	532	15.77	2.39
15.21–17.12	1,942	5.86	687	20.37	3.48
17.12–19.02	1,567	4.73	562	16.66	3.52



Table 4. Continued

Slope Gradient (degree)-1					
Classes	Number of Pixels [ $N_{pix(N)}$ ]	% of $N_{pix(N)}$	Landslide Pixels [ $N_{pix(S)}$ ]	% of $N_{pix(S)}$	FR/PCRV
Geology (lithological composition)-6					
Darjeeling gneiss	6,695	20.21	692	20.52	1.02
Chungtung formation	4,203	12.69	525	15.56	1.23
Lingse granite	3,150	9.51	475	14.08	1.48
Gorubathan formation	2,945	8.89	448	13.28	1.49
Reyang formation	5,925	17.89	621	18.41	1.03
Damuda formation	3,203	9.67	490	14.53	1.50
(Gondwana)					
Siwalik groups	7,010	21.56	122	3.62	0.17
Land Use and Land Cover (LULC)-7					
Tea	2,310	6.97	290	8.60	1.23
Jungle	2,657	8.02	312	9.25	1.15
Open forest	531	1.60	19	0.56	0.35
Degraded forest	1,522	4.59	112	3.32	0.72
Dense forest	2,114	6.38	194	5.75	0.90
Bare surface	4,758	14.36	379	11.24	0.78
Road	1,074	3.24	216	6.40	1.98
Settlement	3,037	9.17	352	10.44	1.14
Agricultural land	6,880	20.77	566	16.78	0.81
Mixed forest	9,281	28.01	933	27.66	0.99
Upslope Contributing Area (UCA, km <sup>2</sup> )-8					
<5.00	11,421	34.47	1089	32.29	0.94
5.00–10.00	7,520	22.70	923	27.36	1.21
10.00–15.00	6,611	19.95	993	29.44	1.48
15.00–20.00	5,215	15.74	253	7.50	0.48
>20.00	2,364	7.14	106	3.14	0.44
Road Contributing Area (RCA, km <sup>2</sup> )-9					
< 0.002	5,720	17.26	0	0	0
0.002–0.004	5,307	16.02	79	2.34	0.15
0.004–0.006	4,220	12.74	440	13.04	1.02
0.006–0.008	4,370	13.19	461	13.67	1.04
0.008–0.010	4,003	12.08	527	15.62	1.29
0.010–0.012	3,522	10.63	572	16.96	1.60
0.012–0.014	2,957	8.93	608	18.03	2.02
>0.014	2,532	7.64	686	30.34	3.97
Settlement Density (No. of settlements/km <sup>2</sup> )-10					
Very low	6,445	19.45	235	6.97	0.36
Low	5,780	17.45	329	9.75	0.56
Moderately low	4,858	14.66	374	11.09	0.76
Moderate	4,397	13.27	499	14.79	1.11
Moderately high	4,265	12.87	591	17.52	1.36
High	3,774	11.39	658	19.50	1.71
Very high	3,612	10.90	687	20.37	1.87

mid-central and mid-lower parts to high (more than 60°) towards the margin and water divide. Most of the landslides occurred in areas with higher than 35° slope gradient, whose FR (prioritized class weight) values range between 1.05 and 1.76. (2) South, southeast, north, east, and northeasterly facing slopes registered highest FR values of 1.29, 1.23, 1.68, 1.53, and 1.15 respectively. All these slope facets were associated with moderate to high landslide susceptibility and a large

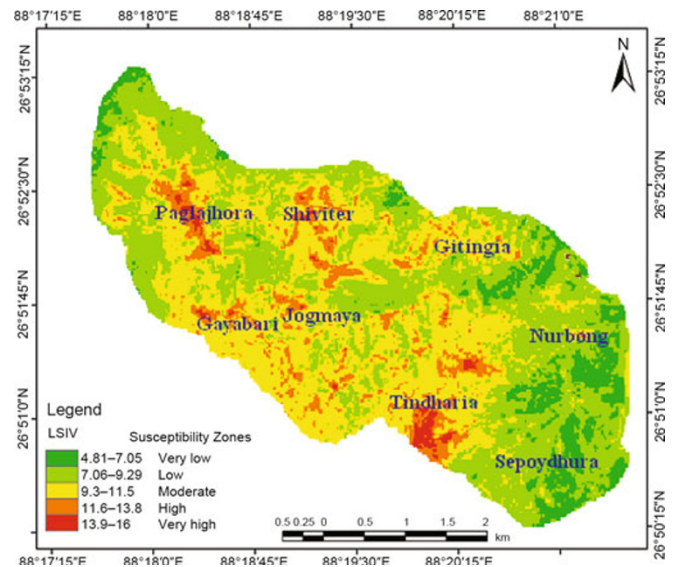
number of landslide occurrences. (3) The derived FR values revealed that high to very high landslide susceptibility zones are characterized by high positive and negative curvature. Lower Paglajhora, Gayabari (Lower), Shiviter (Lower), and Tindharia Tea Estate were characterized by upwardly concave or negative curvature and highest FR values ranging from 1.51 to 2.63. The marginal part of the watershed, mainly Upper Paglajhora, 14 Miles Bustee (upslope), Gayabari

(Upper), and Tindharia (Upper), registered high positive curvature with maximum landslide frequency ratio. (4) Lithologically, darjeeling gneiss, gorubathan, lingste granite, and reyang formations show the highest number of landslide occurrences. Probability of landslide occurrence was very high for the lithological composition of gneiss, mica-schist, and granulite. FR values of lingtse granite, Gorubathan formation, and Chungtung formation were 1.48, 1.49, and 1.23 respectively. All these lithological groups were accompanied with a large number of landslide activities and greater chances of landslide in the future. (5) Drainage density was very high at Lower Paglajhora, Gayabari, and Shiviter Tea Estate, which were characterized by high landslide susceptibility and high FR values ( $>2.5$ ). The value of drainage density increases from the marginal part to the central part. The area with more than 11 km of drainage per km<sup>2</sup> has the highest FR (2.39–3.48) and greater probability of landslide occurrence. (6) The study on lineament showed that most of the major landslide locations are very close to the lineaments. (7) The values of upslope contributing area (UCA) increase from the water divide and the maximum of 20.98 km<sup>2</sup> is registered at the lowermost portion of the watershed. Places with an upslope contributing area of less than 5 km<sup>2</sup> experience less saturation excess run-off and lower intensity of landslides. Larger contributing areas are registered along the main river. This study found that the places with an UCA of 5.00–10.00 km<sup>2</sup> and 10.00–15.00 km<sup>2</sup> have a high FR value of 1.21 and 1.48, which indicates that these places are very prone to landslide hazards. (8) In the Shiv-khola watershed, tea garden, jungle, road, and settlement were characterized by high FR of 1.23, 1.15, 1.98, and 1.14. The analysis shows that tea garden areas, roads, and settlements were dominated by high intensity of landslides and could be treated as maximum probable areas of landslide occurrences. (9) Road contributing area (RCA) is high in Tindharia, Paglajhora, Mahanadi, and Shiviter where the landslide frequency ratio is also very high. At all these places the RCA ranges from 0.008 km<sup>2</sup> to 0.014 km<sup>2</sup> and the prioritized class weight value ranges between 1.04 and 3.97. In the study area, construction of roads and slope modification caused by human intervention are very much responsible for landsliding. (10) The moderate to high density of human settlements at Tindharia, Gayabari, Shiviter, Mahanadi, and Lower Paglajhora are correlated with a high FR as well as greater probability of landslide.

## 4.2 Landslide Susceptibility

In the Shiv-khola watershed, Lower Paglajhora, Shiviter, and Tindharia were very highly susceptible to landslides; Upper Paglajhora, Gayabari, 14 Miles Bustee, and Nurbong Tea Estate were characterized by high landslide susceptibility; Mahanadi and Giddapahar were of moderate landslide potentiality; and the marginal waxing slope of the water divide and the lower-central waning slope areas have low landslide susceptibility (Figure 5).

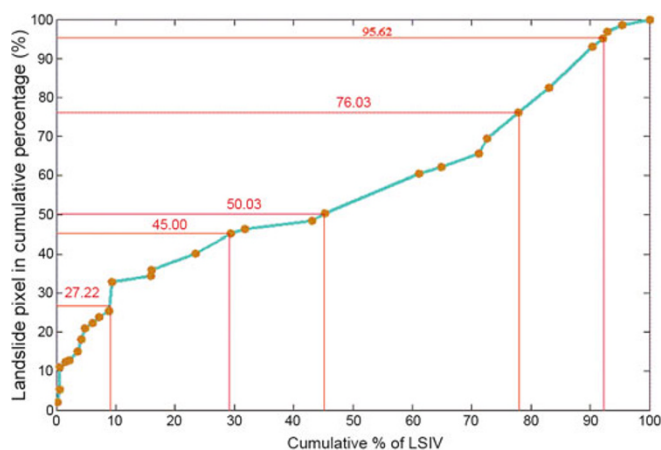
This study revealed that around 50 percent area of the Shiv-khola watershed is in the moderate to very high landslide susceptibility zones with 73 percent of the landslide occurrences. Landslide density in each susceptibility class was derived to evaluate the intensity of landslide activities (Table 5). The landslide density value ranges from 0.03 to 0.25. The calculated density values of 0.25 and 0.15 for very high and high landslide susceptibility zones indicate the higher intensity of landslide activities compared to other landslide susceptibility zones. Landslide density and susceptibility class reveal that the areas with high and very high landslide susceptibility would be prone to fresh landslides and this indicates the validity of the present landslide susceptibility mapping approach.



**Figure 5.** Landslide susceptibility map of Shiv-khola watershed in West Bengal, India

**Table 5.** Relationship between landslide susceptibility, and landslide density

Landslide Susceptibility	Number of Pixels (25 m×25 m) [a]	% of Pixels in Watershed	Landslide Pixels (25 m×25 m)[b]	Landslide Density (b/a)
Very low	7,707	9.03	245	0.0318
Low	35,386	41.46	1247	0.0352
Moderate	34,364	40.26	2676	0.0779
High	6,932	8.12	1074	0.1549
Very high	964	1.30	242	0.2510



**Figure 6. Relationship between landslide affected pixels and landslide susceptibility**

Figure 6 shows the relationship between landslide susceptibility and landslide affected pixels: 27.22, 45.00, 50.03, 76.03, and 95.62 percent landslide affected areas are distributed in 8.75, 28.66, 45, 78, and 92 percent landslide susceptible areas. Around 35 percent landslide affected pixels are distributed in 27 percent of high to very high landslide susceptibility zones, that is, areas with higher probability of landslide activities. But 73 percent landslide susceptible areas contained 65 percent landslide affected pixels.

A comparison between the ground truth data and randomly selected data from the classified image was made on a GIS platform and the result shows that the overall classification accuracy is 92.22 percent, and overall Kappa statistics is 0.894. The class-wise accuracy result is shown in Table 6, which indicates acceptable results.

## 5 Conclusion

This study developed and applied two quantitative analyses that helped to identify landslide susceptible zones in the Shiv-khola watershed. The proposed methodology incorporated all the landslide triggering factors existing in the area. Very fragile and fragmented lithological composition allows

easy percolation of rainwater that generates adequate pore water pressure for promoting downward movement of slope materials in high landslide susceptible sections of the Shiv-khola watershed. The existence of moderate- to high-intensity risk elements and human intervention associated with all favorable geomorphic and geohydrologic landslide triggering factors have made Lower Paglajhora, Tindharia, and Shiviter high to very high landslide hazard risk zones in the Shiv-khola watershed. Slope steepening caused by road-cut benches and toe-erosion, plying of heavy loaded vehicles and their enormous pressure on fragile slope materials, depletion of forest cover at a rapid pace, continuous and regular orographic rainfall in the rainy season, easy percolation of water through fragmented rock-soil composition and increased pore water pressure have caused destructive slope failure, damaged human structures, disrupted normal life by cutting off the communication lines at these three locations, and made these areas the most significant landslide prone sections of Darjeeling Himalaya.

The derived prioritized factor rating values (PFRV) of landslide triggering factors are high for slope gradient (0.2944), lithology (0.2150), drainage (0.1537), and lineament (0.1087), indicating that these are the significant contributing attributes in the Shiv-khola watershed. The PFRV for other risk factors—curvature, upslope contributing area, road contributing area, and land use and land cover, are 0.0764, 0.0535, 0.0375, and 0.0266 respectively, indicating that these factors are also significant landslide risk factors in the study area. The Analytical Hierarchy Process proved to be important to efficiently identify the landslide triggering factors of most importance. The frequency ratio model evaluates the significance of each class of individual factor map in connection to slope instability and their contribution to landslides. These two approaches and their integration are useful for supporting decision making for efficient management. The study shows that geological structure and lineament are important for any structural construction, especially the orientation of roads. Human construction in such areas makes the slopes more risky.

Landslide susceptibility is high in Lower Paglajhora, Tindharia, Shiviter Tea Estate, and Mahanadi Tea Estate, where damages to traffic, life, and property are common

**Table 6. Accuracy assessment: comparison of landslide susceptibility with field data**

Class Name	Classified Total	Number Correct	Producers Correct	Users Accuracy	Accuracy Total
Very Low	0	5	0	0	0
Low	4	3	0	75.00	0
Moderate	11	10	9	90.91	90.00
High	16	15	13	93.75	86.67
Very High	19	17	17	89.47	100.00
Total	50	50	39		

Overall classification accuracy = 92.22%

Overall Kappa statistics = 0.894

phenomena. Plying of heavy loaded vehicles should not be permitted along the Hill Cart Road, no further construction near landslide sites should be allowed, and the expansion of tea estates has to be stopped to reduce the intensity of landslide hazard risks. New laws/regulations should also restrict further construction around the zone of slope failures. Deep-rooted saplings and seeds of grasses should be grown in the landslide prone areas. Land use management should be studied and improved. Development of horizontal and vertical drains to reduce upslope contributing areas and to divert drainage networks from concentrated flow in and around the highly susceptible areas could be the best management options.

## References

- Avinash, K. G., and K. G. Ashamanjari. 2010. A GIS and Frequency Ratio Based Landslide Susceptibility Mapping: Aghnashini River Catchment, Uttara Kannada, India. *International Journal of Geomatics and Geosciences* 1 (3): 343–354.
- Bhattarai, P., and K. Aoyama. 2001. Mass Movement Problems along Prithwi Highway, Nepal. Annual Report of Research Institute for Hazards in Snowy Areas, Niigata University, No. 23, 85–92.
- Borga, M., D. G. Fontana, D. D. Ros, and L. Marchi. 1998. Shallow Landslide Hazard Assessment Using Physically Based Model and Digital Elevation Data. *Environmental Geology* 35 (2–3): 81–88.
- Bradinoni, F., and M. Church. 2004. Representing the Landslide Magnitude Frequency Relation. Capilano River Basin, British Columbia. *Earth Surface Processes and Landforms* 29 (1): 115–124.
- Carson, M. A. 1975. Threshold and Characteristic Angles of Straight Slopes. In: *Proceedings of the 4th Guelph Symposium on Geomorphology*, 19–34. Norwich Geo Books.
- Catlos, E. J., T. M. Harrison, M. J. Kohn, M. Grove, F. J. Ryerson, C. E. Manning, and B. N. Upreti. 2001. Geochronologic and Thermobarometric Constraints on the Evolution on the Main Central Thrust, Central Nepal Himalaya. *Journal of Geophysical Research* 106 (B8): 16177–16204.
- Congalton, R. 1991. A Review of Assessing the Accuracy of Classification of Remotely Sensed Data. *Remote Sensing of Environment* 37 (1): 35–46.
- Dai, F. C., and C. F. Lee. 2002. Landslide Characteristics and Slope Instability Modeling Using GIS, Lantau Island, Hong Kong. *Geomorphology* 42 (3–4): 213–228.
- Dhakal, A. S., T. Amada, and M. Aniya. 2000. Landslide Hazard Mapping and Its Evaluation Using GIS: An Investigation of Sampling Schemes for a Grid-Cell Based Quantitative Method. *Photogrammetric Engineering and Remote Sensing* 66 (8): 981–989.
- Donati, L., and M. C. Turrini. 2002. An Objective and Method to Rank the Importance of the Factors Predisposing to Landslides with the GIS Methodology, Application to an Area of the Apennines (Valnerina; Perugia, Italy). *Engineering Geology* 63 (3–4): 277–289.
- Dutta, K. K. 1966. A Landslip in Darjeeling & Neighbouring Hills Slopes in June. 1950. *Bulletin of the Geological Survey of India*, Series B 15 (1): 7–30.
- Gallant, J. C., and J. P. Wilson. 2000. Primary Topographic Attributes. In: *Terrain Analysis: Principles and Applications*, edited by J. P. Wilson and J. C. Gallant, 51–86. New York: John Wiley & Sons.
- Guzzetti, F., A. Carrara, M. Cardinali, and P. Reichenbach. 1999. Landslide Hazard Evaluation: A Review of Current Techniques and Their Application in a Multi-Scale Study, Central Italy. *Journal of Geomorphology* 31 (1–4): 181–216.
- Intarawichian, N., and S. Dasananda. 2011. Frequency Ratio Model Based Landslide Susceptibility Mapping in Lower Mae Chaem Watershed, Northern Thailand. *Environmental Earth Science* 64 (8): 2271–2285.
- Jadda, M. 2009. Landslide Susceptibility Evaluation and Factor Analysis. *European Journal of Scientific Research* 33 (4): 654–668.
- Jibson, W. R., L. H. Edwin, and A. M. John. 2000. A Method for Producing Digital Probabilistic Seismic Landslide Hazard Maps. *Engineering Geology* 58 (3–4): 271–289.
- Lee, S., and U. Choi. 2003. Development of GIS Based Geological Hazard Information System and Its Application for Landslide Analysis in Korea. *Geoscience Journal* 7 (3): 243–252.
- Lee, S., J. Choi, and K. Min. 2004. Probabilistic Landslide Hazard Mapping Using GIS and Remote Sensing Data at Boun, Korea. *International Journal of Remote Sensing* 25 (11): 2037–2052.
- Lee, S., and B. Pradhan. 2007. Landslide Hazard Mapping at Selangor, Malaysia Using Frequency Ratio and Logistic Regression Models. *Landslides Journal* 4 (1): 33–41.
- Lee, S., J. H. Ryu, J. S. Won, and H. J. Park. 2004. Determination and Publication of the Weights for Landslide Susceptibility Mapping Using an Artificial Neural Network. *Engineering Geology* 71 (3): 289–302.
- Lee, S., and J. A. Talib. 2005. Probabilistic Landslide Susceptibility and Factor Effect Analysis. *Environmental Geology* 47 (7): 982–990.
- Mallet, F. R. 1875. On the Geology and Mineral Resources of the Darjeeling District and Western Duars. *Memoirs of the Geological Survey of India* 2: 1–72.
- Mondal, S., and R. Maiti. 2011. Landslide Susceptibility Analysis of Shiv-khola Watershed, Darjiling: A Remote Sensing & GIS Based Analytical Hierarchy Process (AHP). *Journal of Indian Society of Remote Sensing*. doi:10.1007/s12524-011-0160-9.
- Muthu, K., and M. Petrou. 2007. Landslide Hazard Mapping Using an Expert System and a GIS. *Transactions on Geoscience and Remote Sensing* 45 (2): 522–531.
- Nautilyal, S. P. 1966. On the Stability of Certain Hill Slopes in and Around Darjeeling. *W.B. Bulletin of the Geological Survey of India*, Series B 15 (1): 31–48.
- Pandey, A., P. P. Dabral, V. M. Chowdhary, and N. K. Yadav. 2008. Landslide Hazard Zonation Using Remote Sensing and GIS: A Case Study of Dikrong River Basin, Arunachal Pradesh, India. *Environmental Geology* 54 (7): 1517–1529.
- Pistocchi, A., L. Luzi, and P. Napolitano. 2002. The Use of Predictive Modeling Techniques for Optimal Exploitation of Spatial Databases: A Case Study in Landslide Hazard Mapping with Expert System-Like Methods. *Journal of Environmental Geology* 41 (7): 765–775.
- Pradhan, B. 2010. Remote Sensing and GIS-Based Landslide Hazard Analysis and Cross Validation Using Multivariate Logistic Regression Model on Three Test Areas in Malaysia. *Advance Space Research* 45 (10): 1244–1256.
- Pradhan, B., and S. Lee. 2010a. Delineation of Landslide Hazard Areas on Penang Island, Malaysia, by Using Frequency Ratio, Logistic Regression, and Artificial Neural Network Models. *Environmental Earth Science* 60 (5): 1037–1054.
- Pradhan, B., and S. Lee. 2010b. Regional Landslide Susceptibility Analysis Using Back-Propagation Neural Network Model at Cameron Highland, Malaysia. *Landslides* 7 (1): 13–30.
- Rowbotham, D., and D. N. Dudycha. 1998. GIS Modelling of Slope Stability in Phewa Tal Watershed, Nepal. *Geomorphology* 26 (1–3): 151–170.
- Saaty, T. L. 1977. A Scaling Method for Priorities in Hierarchical Structures. *Journal of Mathematical Psychology* 15 (3): 234–281.
- Saaty, T. L. 1980. *The Analytical Hierarchy Process*. New York: McGraw Hill.
- Saaty, T. L., and L. G. Vargas. 2000. *Models, Methods, Concepts and Applications of the Analytic Hierarchy Process*, 1st Edition. Boston: Kluwer Academic.



- Sarkar, S., and D. P. Kanungo. 2004. An Integrated Approach for Landslide Susceptibility Mapping Using Remote Sensing and GIS. *Photogrammetric Engineering and Remote Sensing* 70 (5): 617–625.
- Sinha-Roy, S. 1982. Himalayan Main Central Thrust and Its Implication for Himalayan Inverted Metamorphism. *Tectonophysics* 84 (2–4): 197–224.
- Tiwari, B., and H. Marui. 2003. Estimation of Residual Shear Strength for Bentonite-Kaolin-Toyoura Sand Mixture. *Journal of Japan Landslide Society* 40 (2): 124–133.
- Tiwari, B., and H. Marui. 2004. Objective Oriented Multi-Stage Ring Shear Test for the Shear Strength of the Landslide Soil. *Journal of Geotechnical and Geoenvironmental Engineering, ASCE* 130 (2): 217–222.
- Vijith, H., and G. Madhu. 2008. Estimating Potential Landslide Sites of an Upland Sub-Watershed in Western Ghat's of Kerala (India) through Frequency Ratio and GIS. *Environmental Geology* 55 (7): 1397–1405.
- Windisch, E. J. 1991. The Hydraulics Problem in Slope Stability Analysis. *Canadian Geotechnical Journal* 28 (6): 903–909.
- Wu, C. Y., and J. P. Qiao. 2009. Relationship Between Landslides and Lithology in the Three Gorges Reservoir Area Based on GIS and Information Value Model. *Frontiers of Forestry in China* 4 (2): 165–170.
- Yalcin, A. 2008. GIS Based Landslide Susceptibility Mapping Using Analytical Hierarchy Process and Bivariate Statistics in Ardesen (Turkey): Comparisons of Results and Confirmations. *Catena* 72 (1): 1–12.
- Yalcin, A., and F. Bulut. 2007. Landslide Susceptibility Mapping Using GIS and Digital Photogrammetric Techniques: A Case Study from Ardesen (NE Turkey). *Journal of Natural Hazard* 41 (1): 201–226.
- Zhou, C. H., C. F. Lee, J. Li, and Z. W. Xu. 2002. On the Spatial Relationship between Landslide and Causative Factors on Lantau Island, Hong Kong. *Geomorphology* 43 (3–4): 197–207.

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